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## Flow Variability and the Biophysical Vitality of River Systems

Robert J. Naiman<sup>3</sup>, Joshua J. Latterell<sup>1</sup>, Neil E. Pettit<sup>2</sup> and Julian D. Olden

School of Aquatic & Fishery Sciences  
University of Washington – 355020  
Seattle, WA 98195 USA

- 1) *Present Address:* King County Department of Natural Resources and Parks  
Water and Land Resources Division  
KSC-NR-0600, 201 S. Jackson St.  
Seattle, WA 98104-3855
- 2) *Present Address:* Centre of Excellence in Natural Resource Management  
University of Western Australia  
Perth 6009, Australia
- 3) *Corresponding Author:* Phone +1 206 685 2025  
Fax +1 206 685 7471  
Email: [naiman@u.washington.edu](mailto:naiman@u.washington.edu)

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## **Abstract**

We illustrate the fundamental importance of fluctuations in natural water flows to the long-term sustainability and productivity of riverine ecosystems and their riparian areas. Natural flows are characterized by temporal and spatial heterogeneity in the magnitude, frequency, duration, timing, rate of change and predictability of discharge. These characteristics, for a specific river or a collection of rivers within a defined region, shape species life histories over evolutionary (millennial) time scales as well as structure the ecological processes and productivity of aquatic and riparian communities. Extreme events – uncommon floods or droughts – are especially important in that they either reset or alter physical and chemical conditions underpinning the long-term development of biotic communities. We present the theoretical rationale for maintaining flow variability to sustain ecological communities and processes, and illustrate the importance of flow variability in two case studies – one from a semi-arid savanna river in South Africa and the other from a temperate rainforest river in North America. We then discuss the scientific challenges of determining the discharge patterns needed for environmental sustainability in a world where rivers, increasingly harnessed for human uses, are experiencing substantially altered flow characteristics relative to their natural states.

## **Introduction**

Extreme natural disturbances – floods, fires, droughts, hurricanes and volcanism – fundamentally alter landscapes, leaving biophysical legacies that persist for centuries. Surprisingly, however, 'extreme' disturbances – though costly to human life and infrastructure – often are not ecologically catastrophic (Turner et al. 1997, 1998). In rivers, for example, extreme floods and droughts are essential for maintaining common biological and physical characteristics, including ecological vitality. We use the term 'vitality' here to reflect the integrity, health, resilience, and productivity underpinning the ability of an ecological system to thrive.

The importance of flow variability is well recognized by ecologists and water resource managers. Nevertheless, water is a valuable commodity as well as a destructive force. Human society requires water for life while seeking protection from floods and droughts. As a consequence, many rivers have been heavily modified to enable water managers to control flows to meet human needs (e.g., industry, agriculture, development of historical floodplains) while dampening or eliminating normal floods and droughts (Gleick 2003). By trapping floods rather than conveying them downstream, many regulated rivers retain little of their original flow variability (Poff et al. 2007). Human control of river flows is now nearly ubiquitous (Vörösmarty et al. 2004, Nilsson et al. 2005). Today, rivers are managed to meet multiple human demands (e.g., steady and dependable water supplies, flood control facilities to protect populated areas). These factors severely constrain the flow variability which is required to meet ecological demands (Naiman et al. 2002, Postal and Richter 2003).

Although human manipulation of river flows provides many societal benefits, it also degrades and eliminates valuable ecosystem services (Richter et al. 2006). River vitality and services (e.g., water purification, recreation) deteriorate when the natural routing of water, sediments and organic materials is substantially modified (Poff et al. 1997, Richter et al. 2003). Remedies to these human-caused problems are often costly or difficult to implement. In the absence of manipulation, floods and droughts interact with the underlying geology to shape the river's physical and chemical templates; sites upon which the biotic communities develop over long time scales (Ward 1998, Lytle and Poff 2004, Naiman et al. 2000, 2005a,b; Figure 1). Significantly modifying or depriving the system of flow extremes upsets patterns in material transport, resource availability, plant mortality and succession, and the composition of biological communities.

Our objective is to illustrate how variability in flow and water temperature shapes the biophysical attributes and functioning of river systems. We explain the ecological rationale for sustaining flow variability. We examine case studies from rivers in two contrasting climate regions – a semi-arid savanna river in South Africa and a temperate rainforest river in North America – that illustrate connections between flow variability, large wood, and the development of river-specific ecological characteristics. We conclude by exploring the importance of variability in establishing environmental flows for rivers – flows needed to sustain ecological systems. This latter subject is especially important in light of ongoing climate changes and the widespread need to balance environmental and human demands for water.

## **The Ecological Rationale for Sustaining Flow Variability**

*Defining flow variability.* The natural flow-regime paradigm postulates that the structure and function of riverine ecosystems, and the adaptations of their constituent riparian and aquatic species, are dictated by patterns of temporal variation in river flows (Poff et al. 1997). In ecological terms, the primary components of a flow regime are the magnitude, frequency, seasonal timing, predictability, duration and rate of change of flow conditions (Table 1). From an evolutionary perspective, extreme events (floods and droughts) exert primary selective pressure for adaptation because they often represent sources of mortality (Poff and Ward 1989, Lytle and Poff 2004).

Substantial spatial variability exists in natural flow characteristics related to climate, geology and topography. Examination of flow gauge data for 806 free-flowing streams and small rivers across the USA by Poff (1996) showed natural flow-regime parameter values ranging from 0.62 to 0.91 for flood frequency (overbank events  $\cdot y^{-1}$ ), 0.44 to 0.96 for flood predictability (proportion of floods falling in a common 60-day period), 2.3 to 9.9 days for average flood duration, and 0.45 to 0.81 for drought predictability (proportion of droughts falling in a common 60-day period). Snowmelt-dominated streams in the Rocky Mountains have the highest predictability of seasonal flood timing ( $\sim 0.95$ ; scale of 0 to 1), relatively high low-flow timing ( $\sim 0.75$ ), and relatively low flood frequency ( $\sim 0.60$ ). By contrast, some streams in the Great Plains of the USA have unpredictable flood ( $\sim 0.45$ ) and low-flow timing ( $\sim 0.50$ ), and very high flood frequency

(~0.90). Subsequent comparisons of river systems across the USA also illustrate substantial within- and between-continent variability in natural flow characteristics (Poff et al. 2007).

*Ecological adaptations to flow regimes.* The physical characteristics of flow variability have strong ecological consequences at local to regional scales, and at time intervals ranging from days (ecological effects) to millennia (evolutionary effects). Even though not all flow regimes favor the evolution of traits that enable flood or drought survival, life-history theory predicts that the magnitude, frequency and predictability of flow events can affect how organisms evolve or fail to evolve (see Lytle and Poff 2004). The theoretical predictions suggest that the timing of life-history events, such as age at reproduction, is optimized with respect to mortality-causing events, such as timing of flood or drought events (Lytle 2001). When extreme flows are frequent, large in magnitude and predictable, selection favors life histories synchronized to avoid or exploit extreme flow events (the evolutionary zone). By contrast, extreme flows that are frequent and large in magnitude but unpredictable have low selection strength for life history timing, even though they might inflict high mortality on populations (the ecological zone). Although no optimal strategy for life-history timing can evolve to avoid unpredictable floods or droughts, bet-hedging strategies might evolve in this case (see Lytle and Poff 2004). For example, in the historically harsh arid environments of the Lower Colorado River basin (USA), evidence suggests that the life-history strategies of native fishes are positioned along an evolutionary bet-hedging axis representing a trade-off between the onset of reproduction and fecundity (Olden et al. 2006). The lifespan of an organism might also affect, and ostensibly reflects, selection strength because disturbance frequency, and thus selection strength, occurs relative to the lifespan of an organism (e.g., a long-lived cottonwood tree experiences annual snowmelt-driven floods many times during a lifetime, whereas a mayfly larvae in the same river system might never experience one).

It has been well-demonstrated that the life history traits of individual species, and the emergent community characteristics, are strongly linked to annual variability in flow regimes (e.g., Poff and Allan 1995, Lamouroux et al. 2002). All parts of the annual hydrograph have consequences for riverine biota (Figure 2). Four principles highlight important mechanisms linking flow variability and aquatic biodiversity, and serve to illustrate the generalized impacts of altered flow regimes (Bunn and Arthington 2002):

- Flow variability is a major determinant of physical habitat in streams, which in turn is a major determinant of biotic composition.
- Many aquatic and riparian species have evolved life history strategies that are responsive to the natural flow regimes.
- Maintenance of natural patterns of longitudinal and lateral connectivity is essential to the viability of populations of riverine species.
- The success of non-native species in rivers is facilitated by the alteration of flow regimes, and the impacts are manifest across broad taxonomic groups including plants, invertebrates, and fish.

*Extreme Variability in Flow – Droughts and Floods.* Uncommon flow levels – droughts and floods – occurring at decadal to century-scale intervals have long-term consequences for biota (Stanley et al. 1997). Whereas most investigations focus on extreme floods (see case studies below), droughts have equally important long-term effects through fundamental changes to biogeochemical processes and to plant and animal distributions (Lake 2003). This is well illustrated in semiarid South Africa where riparian communities are often subjected to multiyear droughts (Naiman et al. 2005b). Normal wet periods are characterized by strong plant production in uplands and in riparian zones, organic litter of good quality, relatively light herbivory by vertebrates with abundant feces returning selective nutrients to the soils, and a general increase in N availability (Figure 3). Excellent upland savanna conditions mean that herbivory is spread across the landscape rather than being concentrated along rivers or in riparian areas. However, savanna conditions deteriorate as drought severity becomes acute. Upland plant biomass and production decline and palatable plant species are over-grazed in riparian corridors. Concomitantly, litter quality and N availability are reduced, the incidence of severe fires in the riparian zone is more frequent, and terrestrial plant species extensively colonize the river corridor. The net result is a '*terrestrialization*' of the riparian zone. It is characterized by a proliferation of successional pathways depending on fire severity, subsequent rainfall, sediment dynamics, and herbivory – and the ecological situation becomes even more complex when rains return (see savanna case study below). Additionally, there are many parallels in rivers with flood-control dams, which cause extended artificial droughts in the historic floodplain.

*Flow Variability and Water Temperature.* Water temperature is a fundamental ecological characteristic – with special significance for maintaining a thriving riverine biota – that is tightly linked to flow patterns. Water temperature is defined by the amount of heat energy entering a stream (which is primarily influenced by climate, channel morphology, groundwater/hyporheic influences, and riparian canopy condition) and the volume of water flowing in the stream channel (Poole and Berman 2001). Therefore, variability in flow regimes affects stream thermal regimes by influencing the timing or magnitude of heat energy delivered to the stream channel (i.e., heat load) or modifying the amount of water (i.e., discharge) that influences the assimilative heat capacity of a stream.

The ecological significance of water temperature in riverine ecosystems is widely recognized. Water temperature directly influences metabolic rates, overall physiology, and life-history traits of aquatic species and helps determine rates of important ecological processes such as nutrient cycling and productivity (Magnuson et al. 1979). Most freshwater organisms are ectotherms, utilizing a diverse array of thermal habitats to meet their specific temperature requirements for survival, growth and reproduction. They respond to the entire thermal regime, which is a composite of patterns of absolute temperatures, diel and seasonal amplitudes, and rates of change; all superimposed upon other related environmental components such as photoperiod (Ward and Stanford 1982). Temporal predictability of a river's thermal regime influences many population parameters of aquatic insects, including egg and larval diapause, embryonic development, larval growth and metabolism (Vannote and Sweeney 1980). The natural thermal regime

of a river also provides temperature cues stimulating fish spawning behavior and reproduction as well as insect emergence and egg hatching.

During drought, water temperatures depend on degree of insolation, substrate type, ground- or rainwater inflows, wind and vegetation cover (Lake 2003). Resistance of insect and fish populations to drought disturbance hinges, in large part, on adaptations of species and the spatiotemporal arrangement of refugia from extremes in temperatures and in dissolved oxygen (Boulton 2003, Matthews and Marsh-Matthews 2003). During extended drought, the distribution of refugia and the frequency and timing of severe temperature episodes play a crucial role in structuring biological communities. Extreme temperatures alter fish physiology and behavior, change competitive interactions, pathogen resistance, and even result in death. Nevertheless, over long time scales, stream populations and communities should be resilient with most species recolonising rewetted streams in days to years, depending upon physiochemical tolerances, distance to source of colonists and extent of drying (Larimore et al. 1959). In other words, the distribution of the population must exceed the spatial extent of the drought or they must persist in refugia. In essence, communities living in systems with recurrent drought are generally adapted to recover relatively quickly from extreme temperature excursions (Poff and Ward 1989).

### **Flow Variability and System Vitality: Case Studies**

#### ***Floods, droughts and the origins of large wood in a semi-arid river***

The flow regimes of arid and semi-arid river systems are highly variable because climatic variability, chiefly rainfall or lack thereof, leads to episodic flooding and prolonged drought (Walker et al. 1995, Rogers and O’Keeffe 2003). For example, the Sabie River, South Africa, experienced a large flood (approximately 100-yr return interval) in February 2000 (Heritage et al. 2001) during which most riparian trees and shrubs were toppled or destroyed (Figure 4). This created approximately 200,000 wood piles along a 105 km stretch of river within Kruger National Park (Pettit et al. 2005). Piles occupied 2-11% of the riverine landscape (Parsons et al. 2006). Wood debris was scarce or absent from most of the riparian corridor prior to the flood (Parsons et al. 2006). Subsequently, it became apparent that the flood was a primary event driving the long-term ecological characteristics through the conversion of the riparian trees to woody material in the river corridor. While infrequent, events similar to this appear to shape the structure of riparian forests for decades to centuries (Turner et al. 1998).

The wood deposited in the Sabie River after the 2000 flood reflected recent terrestrialization of the riparian zone due to a decadal long drought, which allowed terrestrial trees to colonize the channel floor (van Coller et al. 1997, Rogers and O’Keeffe 2003). A previous flood of comparable size on the Sabie River occurred in 1925 (Heritage et al. 2001) and there is some evidence that flood established conditions for a significant recruitment of some important riparian trees. Prior to the 2000 flood the age-class distribution of the large riparian tree *Ficus sycamorus* (Sycamore fig) indicated that there had been little recruitment in recent years and possibly a large episodic recruitment

event occurred approximately 70 years previous (van der Velde 2001). This suggests that many of the mature trees present in 2000 were likely to have recruited from the 1925 flood. Wood piles are important for the regeneration of riparian trees such as *F. sycamorus*, as seedlings of this species establish preferentially in wood piles after floods (Figure 5; Pettit and Naiman 2006).

*Influence of large wood (LW) on initial successional pathways.* Conditions created by LW debris in the post-flood environment suggest that type, abundance and position of wood piles shape the overall pathways of vegetation succession. Initial establishment from seed after the 2000 flood was dominated by ruderals (particularly grasses and annual forbs) in the channel and riparian areas (Pettit and Naiman 2005). However woody plant regeneration (in terms of number of species and cover) within piles was significantly higher than outside piles (Table 2). This was due mainly to resprouting of toppled trees with intact roots and living tissues (36% of LW surveyed), although there was some regeneration from seedlings. This allows for the rapid re-establishment of vegetative biomass in the early stages of post-flood riparian forest renewal, and nearly all are associated with LW piles.

Deposited LW acts as localized focal points for accumulating fine sediments, soil nutrients and soil moisture that, collectively, results in patchy resource availability (Pettit and Naiman 2005). Favorable microclimates (e.g., increased soil moisture, reduced summer soil temperatures) and soil nutrients (e.g., elevated levels of N, P, K) within LW piles remain evident for years afterward (Table 2). These beneficial and stable conditions within LW piles make them particularly important as refugia for plants in semi-arid environments with long dry seasons, high ambient temperatures and considerable climatic variability. Even slight enhancement of microclimates and nutrient availability significantly improves the survival of germinating seeds and developing seedlings, particularly during stressful periods of low moisture and high temperature.

Large wood piles are also fuel sources that may allow seasonal savanna fires to penetrate the riparian zone; thereby elevating the risk of high severity fires, greater fire frequency and longer fire residence times. Dead wood accumulated around standing trees creates a particularly acute fire hazard for living trees, even in relatively mild fires. Patchiness of the riparian landscape (heterogeneity) is increased by fire and the development of resource-rich patches is enhanced by the presence of LW piles within the burnt landscape (Figure 6; Pettit and Naiman 2007a,b). This is seen in patterns of soil nutrient distribution (soil P 49% higher in burnt wood patches after 24 months) and vegetation cover with vegetation sparse in intensively burnt areas such as where LW was deposited (58% reduction in burnt wood patches). Fire patchiness also enables the continued existence of fire-sensitive species in riparian zones by providing unburnt refugia.

*Large wood and riparian heterogeneity in semi-arid rivers.* LW creates focal points for regeneration after large infrequent floods. These focal points provide a blueprint for patterns of regeneration as well as contribute to the spatial heterogeneity and biocomplexity (Figure 7, Pettit and Naiman 2005). Despite the relatively brief retention time, the ecological legacy of LW deposition in the Sabie River appears to be preserved

for long periods, manifested in vegetation structure, depositional features, buried wood deposits and local fire impacts on surviving vegetation. In arid and semi-arid environments, the patchy environment is effective at trapping and retaining resources; thus using them efficiently (Ludwig and Tongway 1995). The dynamics and spatial variability of material inputs (including LW) also contribute to heterogeneity and hence the resilience of riparian landscapes.

***Flow variability and the ecological characteristics of a temperate rainforest river***

Flow variability supports temperate floodplain rivers in a manner similar to, but with slightly different mechanisms, from the semi-arid river described above. High flows capture LW by eroding stream banks and causing trees to fall into the channel; the resulting LW plays an important role equal to that in semi-arid rivers in shaping aquatic and riparian habitats. The ecological roles of LW in temperate regions are better quantified than in semi-arid regions but the functional roles are similar (Gregory et al. 2003, Pettit et al. 2006). In-channel LW retains organic matter and moisture, forms pools, and promotes nutrient uptake (Valett et al. 2002). Remnant logs function as habitat for terrestrial organisms and young plants. Many remnant logs contained in early-seral floodplain forests eventually re-enter the channel through ongoing bank erosion and continue to move downstream, creating habitat in numerous locations before disintegrating (Latterell and Naiman 2007). Some remain stable for decades or more and protect downstream forests from erosion for long periods (Montgomery and Abbe 2006). The complexity of the river-floodplain system is supported by strong feedbacks and interactions between wood, pioneering vegetation, and sediments – with flow variability acting as a master variable.

Interactions between wood dynamics and flow magnitude have been widely demonstrated (e.g., Piégay 2003, van der Nat et al. 2003a,b) and are commonly observed during floods. Though floods are ecologically important events, the physical characteristics of aquatic and riparian habitats in floodplain rivers reflect the consequences of the entire flow regime (Poff et al. 1997). For example, channel migration in the Yampa River, Colorado (USA) is linked to the duration of flooding at 125% of bankfull discharge rather than simply the absolute flood magnitude (Richter and Richter 2000). In many other systems, such as the semi-arid river described above, large but infrequent floods erode riparian forests (Piégay and Bravard 1997, Parsons et al. 2005), generating massive quantities of LW (Pettit et al. 2005; Pettit and Naiman 2005). In general, most lines of evidence indicate that high flows are important for recruiting, mobilizing and transporting wood in floodplain rivers (Marcus et al. 2002) but transport also occurs during lesser flows (van der Nat et al. 2003a). Flow variability and resulting channel dynamics cause wood to 'spiral' downstream (Latterell and Naiman 2007). Wood, much like other particles, moves at irregular intervals that are, in part, flow dependent. At other times LW deposits at specific places in the channel until the return of necessary hydraulic conditions for movement. Instream supplies of wood therefore consist of a mixture of logs recently recruited to the channel from the surrounding forest as well as some from the distant past (e.g., Hyatt and Naiman 2001). The transient storage of LW in the floodplain likely buffers against climatic periods when delivery rates of new wood are low.

Observations from the Queets River in Olympic National Park, Washington USA, provide corroborating evidence that flow variability influences wood recruitment. Unlike the Sabie River, short duration catastrophic floods have not been observed on the Queets River, perhaps because of the relatively short period of record. However, rainfall-driven flow variability is substantial within seasons and among years. Reconstructions from sequences of historical air photos suggest that temporal variation in input rates of large trees from the erosion of mature terrace forests is positively related to the average magnitude of annual peak flow occurring in a given period of time ( $r^2 = 0.92$ ;  $p = 0.04$ ; Figure 8). Most LW originates where channel meanders undercut Sitka spruce (*Picea sitchensis*) and western hemlock (*Tsuga heterophylla*) in mature river terraces flanking the downstream edges of meander bends (Latterell and Naiman 2007).

Wood inputs resulting from flow variability aid in the initiation of a patchwork of forest regeneration on alluvial surfaces. As in the semi-arid rivers, successional processes reorganize the riparian plant community, gradually transforming bare alluvial deposits into forests over decades to centuries (Décamps 1996). In the Queets River, pioneering vegetation is estimated to colonize over 95% of the channel area in a given year within a century – but often unsuccessfully as a result of ongoing flooding and channel movement (Figure 9; Latterell et al. 2006). Flow variability interacts with the heterogeneous topography of the floodplain and the roughness created by LW and pioneering vegetation, influencing soil depth and stand structure. Ongoing interactions, combined with differences in site history and the time elapsed since patches were formed, cause patches to have contrasting vegetation age and size distribution, species composition, microclimate, soil characteristics and LW deposits (Naiman et al. 2005a,b; Latterell et al. 2006, Van Pelt et al. 2006). Eventually, many patches are destroyed by channel movements. In the Queets River, half of the floodplain forest area is eroded within two decades after being formed, though four centuries are estimated to pass before half of the mature fluvial terraces is eroded (Latterell et al. 2006). Destruction of forest patches provides LW, sediment, and propagules (seeds, roots, and branch fragments) that initiate new patches downstream, where floodplain development begins again.

Ongoing patch destruction and periodic rearrangement of human-placed LW by floods is sometimes viewed as an undesirable outcome from the point of river restoration. This is understandable when engineered structures are intended to provide bank protection, or the river system is so constrained by human pressures that insufficient resources exist upstream to replace LW deposits and forest patches destroyed by the river. However, we argue the natural destruction and creation of some floodplain patches is actually needed to sustain the dynamic patch mosaic and may contribute to system resilience by stabilizing the availability of certain habitats at large spatial scales (e.g., Arscott et al. 2002, van der Nat et al. 2003b, Gurnell et al. 2005, Latterell et al. 2006). For example, the distribution of pools (Naiman and Latterell 2005) and pioneer bars (Latterell et al. 2006) changes substantially over time in the Queets River, as patches are destroyed and re-created. These changes are fundamental characteristics of the system; the distribution of individual channel units or habitats fluctuates over time but the patterns of change are asynchronous among reaches, dampening the degree of change at the valley scale. As a

result, they contribute to the resilience of riverine species that utilize these habitats for parts of their life cycle.

### **The Challenge Ahead: Establishing Environmental Flows**

The act of explicitly managing water flows for rivers – environmental flows – allows an appropriate volume and timing of water flows to support the ecological health and the livelihoods of dependent organisms, and in an ideal world this would include appropriate flow extremes (Figure 10). The case studies illustrate the important influences of flow variability on fundamental, long-term ecological processes in natural rivers. A fundamental challenge for managers is to balance human and environmental needs for water (Naiman et al. 2002). Hundreds of approaches have been tried in the past few decades with varying degrees of success with respect to flow extremes (Tharme 2003). At first blush, creative solutions abound. In practice, entrenched human demands on rivers create nearly insurmountable obstacles to achieving true balance. Establishing environmental flows promises to be effective because the imbalance is rectified by strategically re-aligning key human flow alterations rather than removing humans from the equation. In this way, the water management community may succeed in improving river vitality and permit society to enjoy the ecological goods and services that healthy rivers provide.

The primary challenge to balancing human and environmental needs for water is that flow alterations, channel modifications, and land use changes are often inseparably linked, particularly in large rivers where flow modification is most pervasive. This practical constraint is unavoidable in many human-modified landscapes (e.g., agroecosystems to urbanized and industrial). Successful flow rehabilitation will often require simultaneous efforts to restore channel form and reconnect riparian areas, including floodplains, which in turn requires transformation of existing land uses. Land use transformation will require great care to avoid causing irreversible social hardship. Existing human demands of many large rivers are so great that striking a balance with the needs of the environment require a level of technical effort and social investment that is as yet unprecedented.

Knowledge gaps also present significant challenges to recovering flow variability in modified rivers. Natural differences in flow variability among rivers needs to be better quantified. Empirical relationships between flow variation and ecological responses (e.g., protection of freshwater biodiversity and the maintenance of goods and services that rivers provide) remain elusive. Solving these knowledge gaps remains a great challenge for water managers and scientists alike (Arthington et al. 2006). The technical challenges and the practicalities of planning or implementing extreme flow events are not trivial.

Despite considerable progress in understanding how flow variability sustains river ecosystems there is, nevertheless, a growing temptation to ignore natural system complexity – especially the importance of flow variability – in favor of simplistic, static, environmental flow “rules” to resolve pressing freshwater issues. Such approaches are misguided and ultimately fail to sustain vitality in rivers. Water managers everywhere are

increasingly required to provide reliable and affordable water supplies to a growing population while, at the same time, are expected to do so without degrading freshwater ecosystems (Tharme 2003, Postel and Richter 2003, Richter et al. 2003). Fortunately, this challenge is recognized, and many scientists and managers are working closely to develop a process (described below) that can be implemented at regional scales (Arthington et al. 2006, Poff et al. 2008).

Determining and implementing effective environmental flow regimes for rivers acknowledges the need for appropriate variability in discharge (Arthington et al. 2006, Poff et al. 2008) and temperatures (Olden and Naiman 2008). Defining 'appropriate variability' will be increasingly difficult, because climate change renders the past an imperfect analogue of future conditions. Environmental flow guidelines must therefore explicitly incorporate likely consequences of climate change on the hydrology of managed systems to reduce the likelihood of future crises. Sustaining healthy freshwater ecosystems requires that adequate water flows – and associated temperatures – are maintained within the ecosystem while other human uses of water (e.g., urban, agricultural, power generation, flood control) are accommodated.

Generic approaches for determining environmental flow requirements are being developed. The most promising ones incorporate essential aspects of natural flow variability shared across particular classes of rivers that can be validated with empirical biological data and other information in a calibration process (Arthington et al. 2006). One approach under development is the "Ecological Limits of Hydrologic Alteration," or ELOHA (Poff et al. 2008). This method appears to satisfy several essential scientific requirements for successful application even though the social and political dimensions require further development. In addition to being grounded in mechanistically-based, flow-ecology linkages that are subject to empirical testing and validation, it is applicable

- at a regional scale
- across a spectrum of worldwide social, political and governance contexts, and useful regardless of the stage of water resource development and historical status of environmental flow protection
- across an array of flow alterations, from modified land use to river regulation by dams
- across a wide range of available data and scientific capacities

There are four steps general scientific steps in ELOHA (Poff et al. 2008). They are: (1) building a "hydrologic foundation" of streamflow time series for both baseline (undeveloped) and developed conditions, throughout the region of interest; (2) classifying rivers throughout the region into "river types," using hydrologic time series that represent baseline conditions; (3) estimating the degree of hydrologic alteration that has occurred throughout the region, based on comparison between baseline and developed conditions, resulting in measures of hydrologic alteration; and, (4) developing flow-ecology response curves for each river type by associating increasing degrees of hydrologic alteration with increasing or decreasing ecological change.

The need for practical approaches to environmental flows, like ELOHA, is global. Despite hundreds of international treaties, conventions, and national and state policies and laws for ecosystem protection, the global degradation of freshwater biodiversity and environmental quality is ongoing, even accelerating (MEA 2005, Dudgeon et al. 2006). Much of this degradation is a direct result of flow homogenization of the world's rivers by dams and by water withdrawals that undermine natural flow variability (Gleick 2003, Vörösmarty et al. 2004, Poff et al. 2007). Nevertheless, it is recognized that flow regulation, land fragmentation and development are a suite of tightly interacting factors, often implemented simultaneously, making it difficult to assign cause and effect to one or the other. The issue becomes even more complex when one considers the ongoing and widespread changes to flow regimes from climate change (Battin et al. 2007, IPCC 2007) and the enormous efforts expended in river restoration that largely ignore the fundamental importance of flow variability (Bernhardt et al. 2005).

Sustainable water resource management is constrained by three pervasive myths; that societal and environmental water demands always compete with one another; that technological solutions can solve all water resource management problems; and that environmental solutions to protect and maintain freshwater resources are more expensive and less dependable than technological solutions (Bernhardt et al. 2006). While conservation and good stewardship of water resources can go a long way toward meeting societal demands and values, new approaches to sustain ecosystem health and biodiversity in rivers and their associated systems can be well aligned with options for human use. These options can deliver a suite of ecosystem goods and services to society – but only if ecologically-appropriate flow variability is maintained.

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Table 1. Common terms used to describe variability in river flows.

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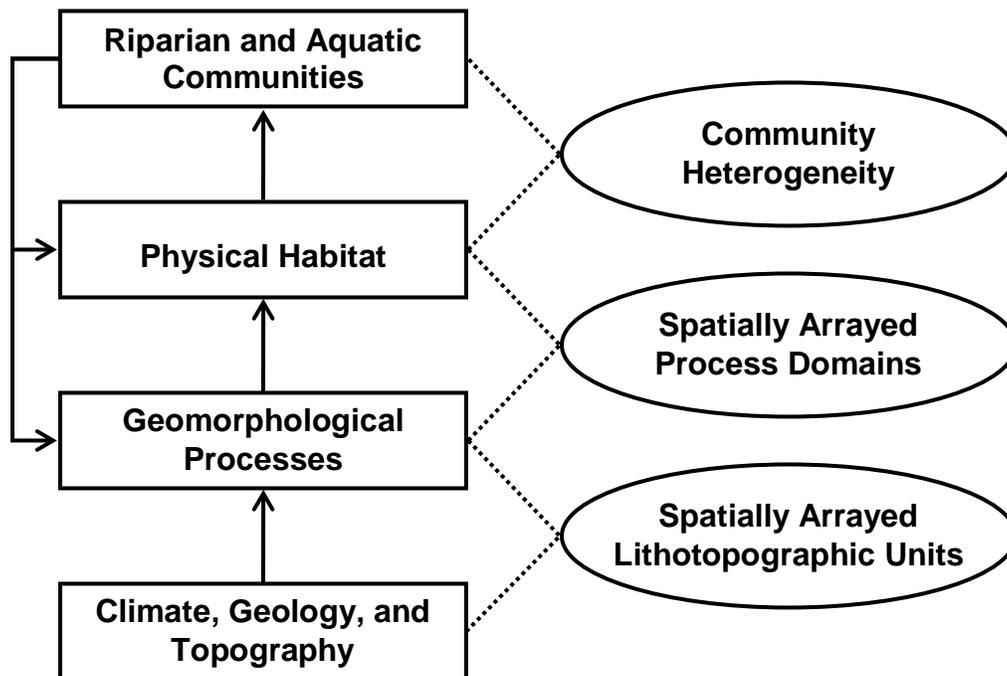
- Magnitude: the amount of water moving past a fixed location per unit time. The larger (or smaller) the magnitude of a flood (or drought), the greater the expected physical impact.
  - Frequency: the number of events of a given magnitude per time interval (e.g., per year). For a given river or stream, frequency is typically related inversely to magnitude.
  - Duration: the period of time associated with a particular flow event. Typically expressed in terms of number of days a flood or drought lasts.
  - Timing: the date during the year that flood or drought occurs, often derived from long-term flow records.
  - Predictability: the degree to which flood or drought events are autocorrelated temporally, typically on an annual cycle. Predictable events also might be correlated with other environmental signals (e.g., rainfall events, seasonal thermal extremes, photoperiod, sudden increases or decreases in flow).
-

**Table 2.** Mean values for LWD piles and adjacent reference (non wood) plots in soil nutrient, soil physical properties and vegetation across all landscape positions and LWD pile types. Values are mean  $\pm$  standard error from a sample size of 119 paired LWD and reference plots.

	LWD pile	Reference plot	t calculated <sup>1</sup>	p <sup>1</sup>
<i>Soil nutrients</i>				
Total N (%)	0.097 $\pm$ 0.001	0.081 $\pm$ 0.001	3.59	0.005
Available P ( $\mu$ g/g)	10.4 $\pm$ 0.7	6.9 $\pm$ 0.54		<0.001
K (mg/l)	196.1 $\pm$ 11.7	122.5 $\pm$ 8.3		<0.001
Total C (%)	0.87 $\pm$ 0.04	0.64 $\pm$ 0.05	3.91	<0.001
C:N ratio	8.69 $\pm$ 0.89	7.11 $\pm$ 0.56	6.24	<0.001
<i>Soil physical</i>				
Silt & clay (%)	22.5 $\pm$ 1.2	16.2 $\pm$ 1.28	5.87	<0.001
Temperature (C)	27.7 $\pm$ 0.21	33.4 $\pm$ 1.41		<0.001
Soil Moisture (%)	10.6 $\pm$ 0.32	5.7 $\pm$ 0.43	9.18	<0.001
<i>Vegetation</i>				
Foliage cover (%)	62.5 $\pm$ 8.2	50.7 $\pm$ 4.7	2.28	0.025
No. species	11.4 $\pm$ 0.4	10.8 $\pm$ 0.41	1.39	0.17
Cover woody spp (%)	33.2 $\pm$ 17	9.2 $\pm$ 6	6.33	<0.001
No. woody spp	3.8 $\pm$ 0.5	2.6 $\pm$ 0.4	5.41	<0.001
<sup>2</sup> No. tree seedlings (ha <sup>-1</sup> )	19.9 $\pm$ 5	16.6 $\pm$ 4	1.89	0.12

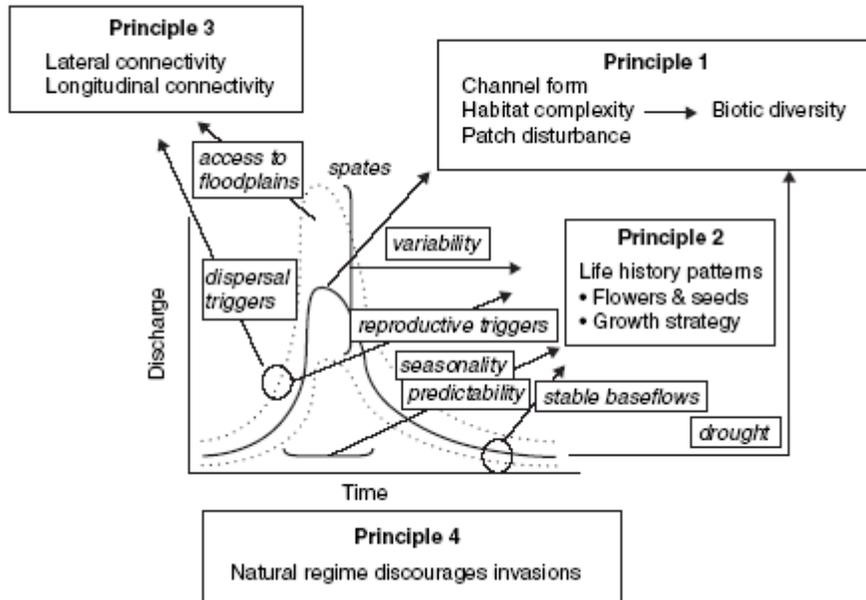
<sup>1</sup> paired t test

<sup>2</sup> Mean proportion of area for each substrate was wood pile = 10.25  $\pm$  4.6%; non wood (open or vegetation) = 65.4  $\pm$  6.2%.

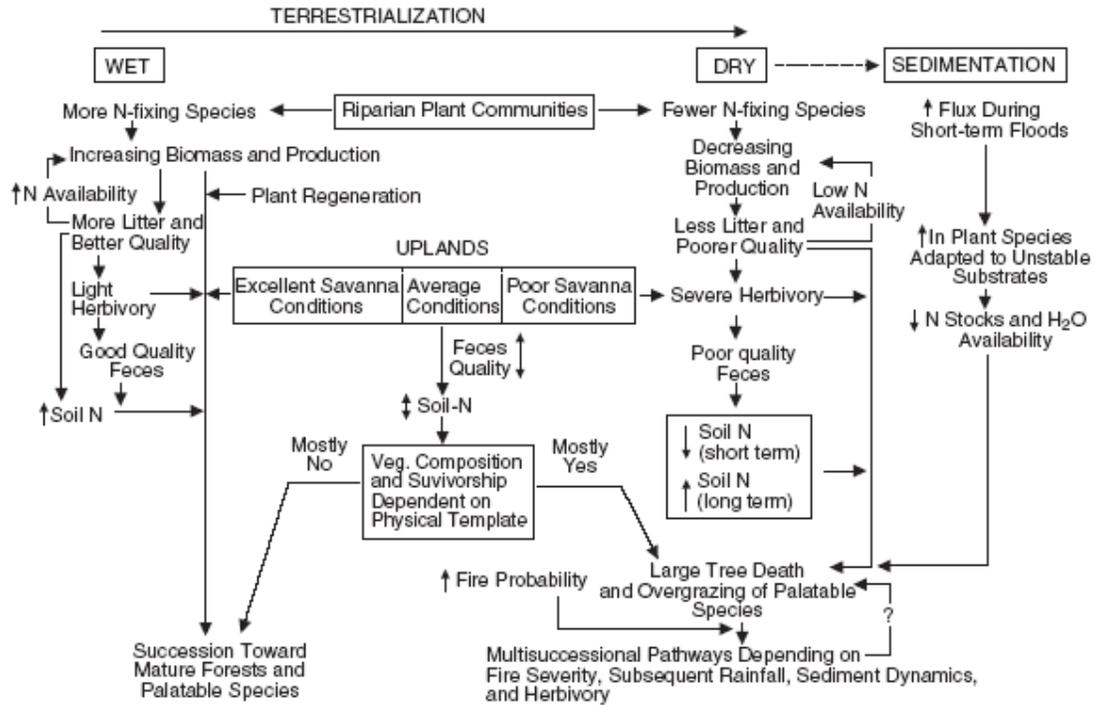


**Figure 1.** Schematic representation of the relationship among hydrogeomorphic processes, habitat dynamics, and riverine communities. Lithotopographic units are areas with similar topography and geology, and within which similar suites of geomorphic processes occur. Modified from Montgomery (1999) and from Naiman et al. (2000).

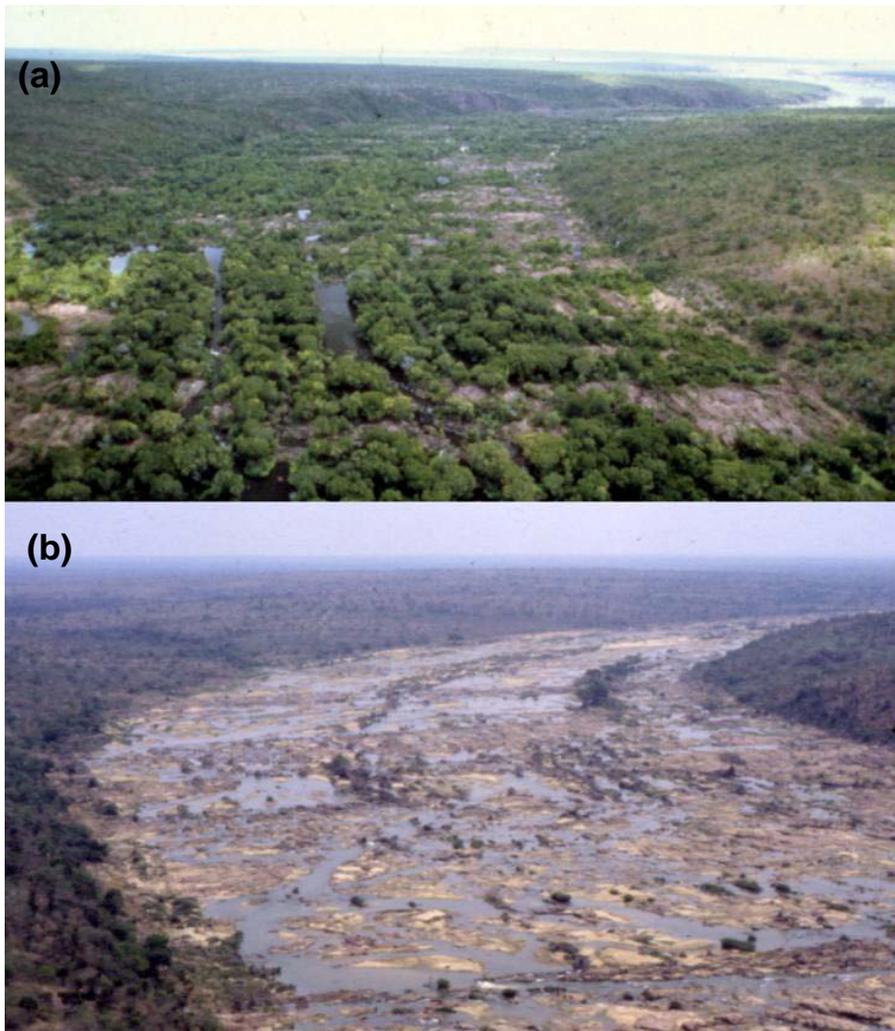
## Riparian Biodiversity and Natural Flow Regimes



**Figure 2.** The natural flow regime influences riparian biodiversity via several interrelated mechanisms operating over different spatial and temporal scales. The relationship between biodiversity and the physical nature of the riparian habitat is driven primarily by large floods that shape channel form (Principle 1). However, droughts and base-flows also play roles by limiting overall habitat availability. Many features of the flow regime influence life history patterns, especially the seasonality and predictability of the pattern (Principle 2). Some flows trigger longitudinal dispersal of riparian organisms and other large events inundate otherwise disconnected floodplain habitats (Principle 3). The native biota have evolved in response to the overall flow regime. Invasion by non-native species are more likely to succeed at the expense of native biota if the former are better adapted to modified flow regimes (Principle 4). Adapted from Bunn and Arthington (2002).



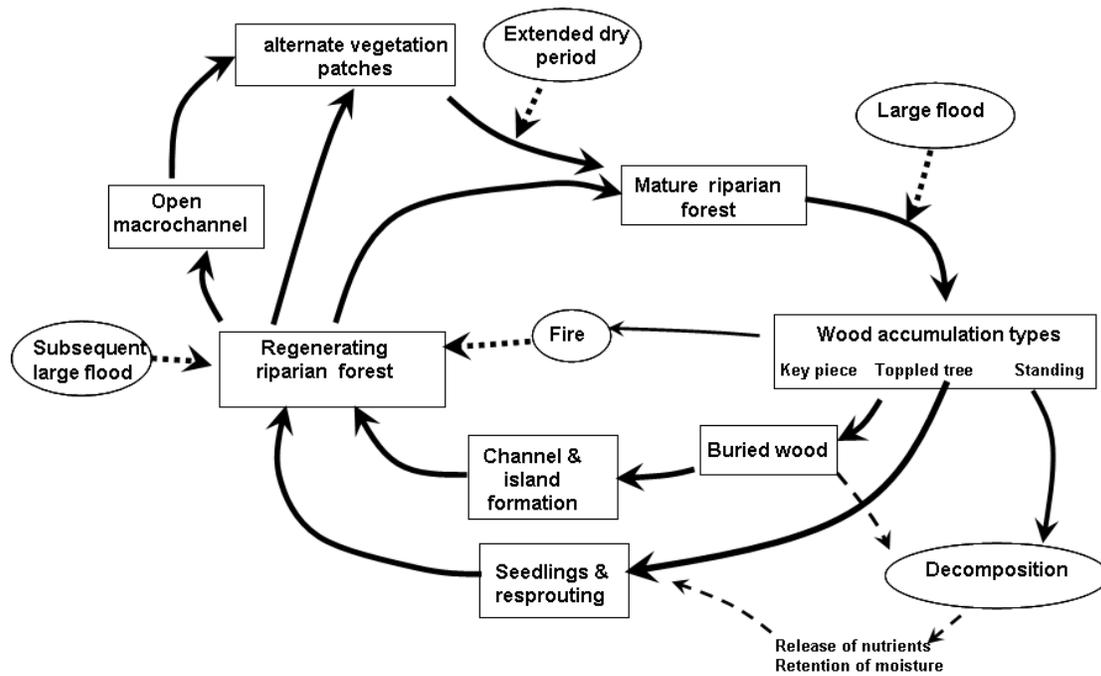
**Figure 3.** A conceptual model of the environmental responses of riparian corridors in the semi-arid region of South Africa to increasing drought conditions. In effect, the riparian system becomes more and more terrestrial as species and processes normally associated with the savanna uplands move closer to the river (from Naiman et al. 2005b)



**Figure 4:** Oblique aerial view of the Sabie River, Kruger National Park. a) Pre-flood where there were extensive areas of riparian forest in the macrochannel. b) Post-flood after a large proportion of the forest had been destroyed. (Photo: F. Venter, Kruger National Park, South Africa)



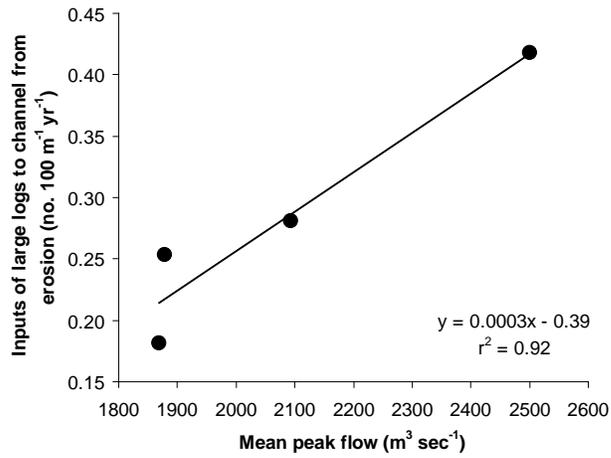
**Figure 5:** Development of tree seedlings in large wood deposited after flooding on the Sabie River in February 2000. a) *Diospyros. mespiliformis* (jackal berry). b) *Ficus sycamorus* (Sycamore fig). (Photos: N. Pettit)



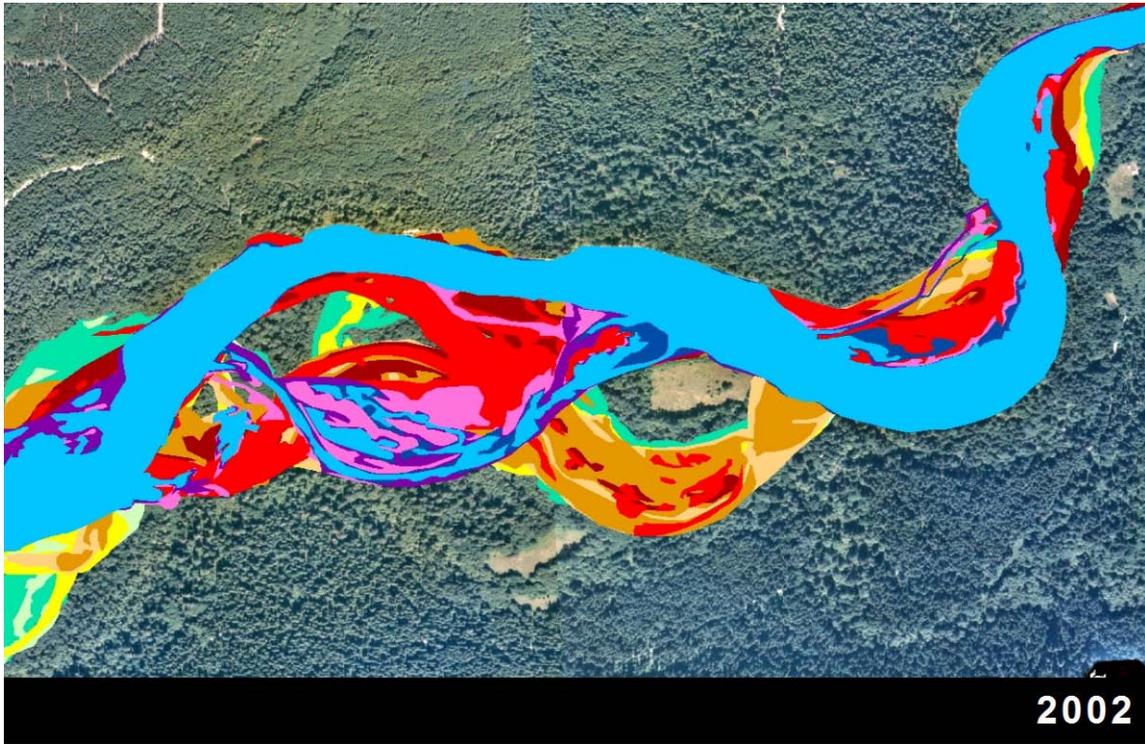
**Figure 6:** Pathways of influence of wood accumulations of successional processes in riparian vegetation and landscape heterogeneity after a large flood on the Sabie River, Kruger National Park, South Africa.



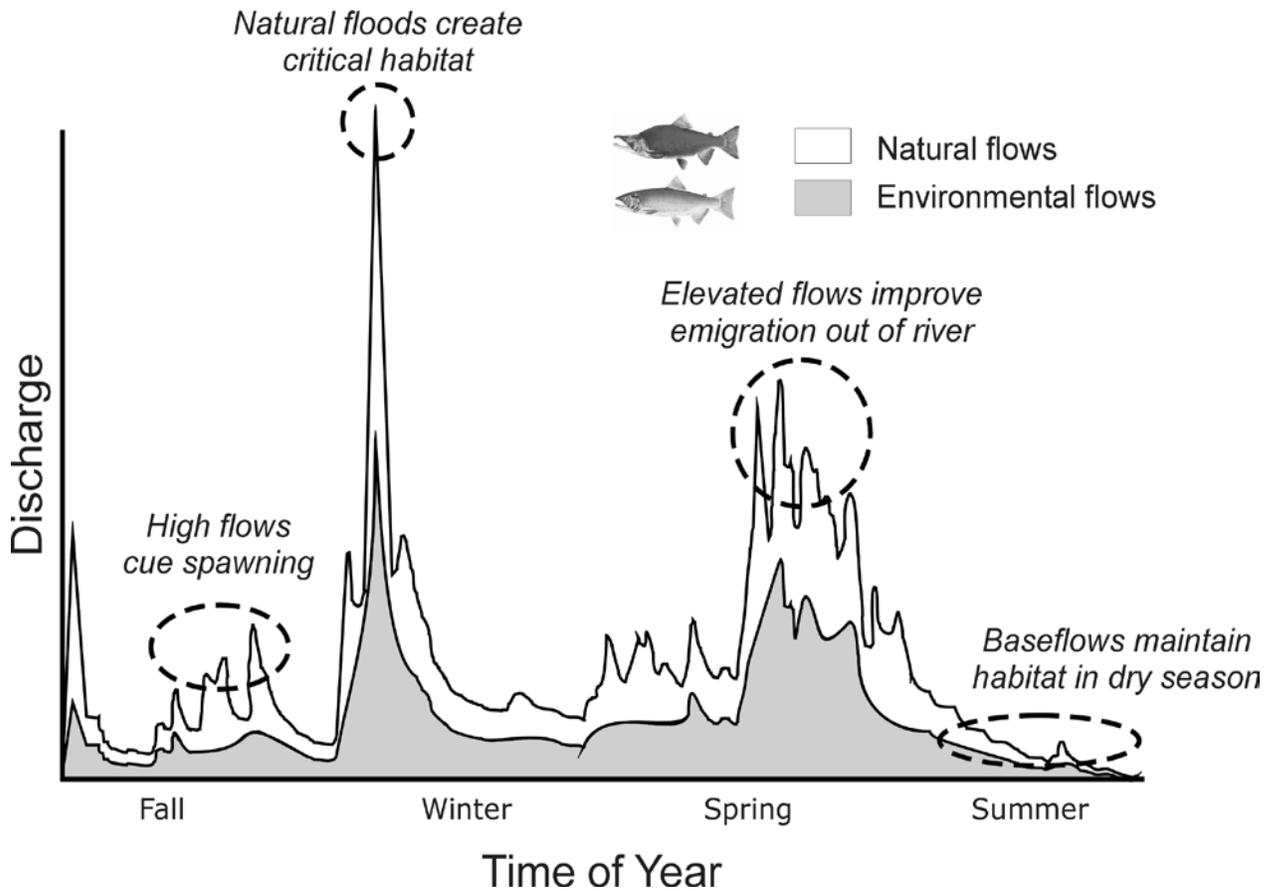
**Figure 7:** Burnt riparian area of the Sabie River. Note the ash bed from a burnt out large wood pile in the foreground, burnt trees resprouting from the base (mid-ground) and unburnt trees in the background. (Photo: R Naiman)



**Figure 8.** Relationship between annual input rate of large logs to the channel from erosion of mature forest terraces and mean peak flow rates over the entire Queets River mainstem in Olympic National Park, Washington, USA. Inputs were estimated from the area of forest eroded over an interval between aerial photographs (see Latterell and Naiman 2007 for details). Mean peak flows were calculated from USGS records over each corresponding interval.



**Figure 9.** Variation in channel position in the Queets River, Washington, USA at 11 different years over a 63-year period: 1939, 1950, 1954, 1962, 1968, 1973, 1981, 1993, 1997, 2000, 2002. The scale of the image is roughly 2.5 km from left to right. (from Van Pelt et al. 2006)



**Figure 10.** Comparative ability of natural and environmental flows to create and maintain U.S. Pacific salmon habitat and cue spawning. Shown is a hypothetical hydrograph that largely mimics temporal discharge patterns of the Trinity River, California (U.S.A.) and the life-history requirements of Chinook salmon (*Oncorhynchus tshawytscha*).